

65015

DRAFT □

Poikilitic Impact Melt Breccia  
1802 grams



Figure 1: Photo of 65015 showing rounded exposed surface with numerous zap pits and angular portion without zap pits protected by burial in soil. NASA photo # S72-43936. Cube and scale are 1 cm.

### **Introduction**

Sample 65015 was collected from the lower slope of Stone Mountain. It was about half buried in the soil, as judged by the obvious “soil line”, below which there are no micrometeorite pits (figure 1-3). It is a very coherent, dense rock with little void space.

The research on 65015 is well summarized by Ryder and Norman (1980).

### **Petrography**

65015 is a feldspathic impact melt characterized by a well developed poikilitic texture in which oikocrysts of pyroxene enclose abundant clasts and chadocrysts (figures 5-7). Most, but not all, of the chadocrysts are plagioclase. Albee et al. (1973) also report numerous relict lithic fragments as clasts in 65015.

Large interlocking oikocrysts of pyroxene (up to 0.6 mm) dominate the texture (figures 6-7). Boundaries between the oikocrysts are characterized by lath-shaped plagioclase intergrown with granular olivine and accessory ilmenite, Fe-Ni metal, troilite, phosphates and K-feldspar. The ilmenite (0.2 mm) occurs in chains and is itself poikilitic. Abundant plagioclase grains (up to 1.5 mm), occurring as chadocrysts throughout the rock, are found to be very calcic ( $An_{93-95}$ ). Papanastassiou and Wasserburg (1973) and Meyer et al. (1974) found that the large plagioclase grains were not in equilibrium with the trace element content of the bulk sample.

The rock has a high content of trace elements and meteoritic siderophiles (Ni, Ir, Au). Wasson et al. (1975) found that the ratios of siderophile elements in metal grains separated from the rock were different from those in the bulk rock.



Figure 2: End view of 65015 showing angular surface that was buried in the soil. NASA # S72-43935

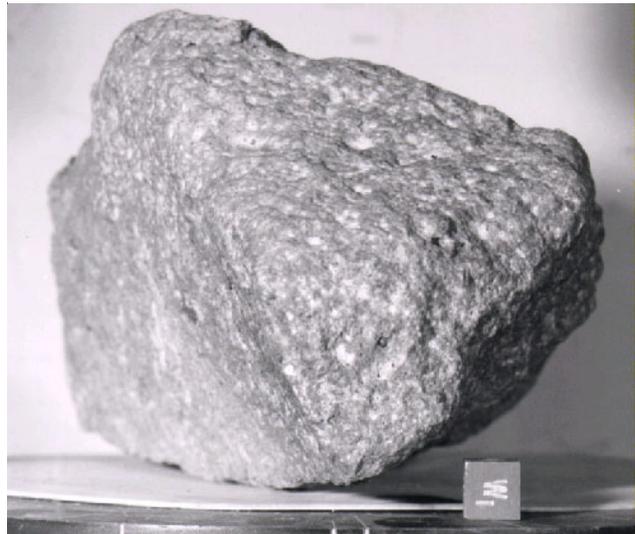


Figure 3: End view of 65015 illustrating rounding caused by micrometeorite bombardment. NASA # S72-43939

### Mineralogical Mode

	Albee et al. 1973	Simonds et al. 1973
Plagioclase	57.1 vol. %	61
Low-Ca pyroxene	28.9	29
High-Ca pyroxene	6.4	6
Olivine	1.1	1
Opaques	1.7	3
Mesostasis	3.6	
Whitlockite	0.7	

The distinctive poikilitic texture of this rock, the inclusion of relict lithic and mineral clasts and the high meteoritic siderophile content are also common features of the Apollo 17, Boulder 6 samples as well as the melt sheets found in large terrestrial craters. Simonds et al. (1976) modeled the formation of these rocks as a mix of impact melt and cold clastic material.

### Mineralogy

**Olivine:** Relict, partially reacted, olivine grains ( $Fo_{65-74}$ ) are reported by Albee et al. (1973)

**Pyroxene:** Albee et al. (1973) determined the chemical composition of pyroxene in 65015 (figure 4). The low-Ca pyroxene occurs as large oikocrysts and is generally homogenous. The high-Ca pyroxene is found in the intergranular regions and in clasts and is somewhat variable in composition.

**Plagioclase:** Abundant plagioclase grains (up to 1.5 mm) occur as chadocrysts throughout the rock and are

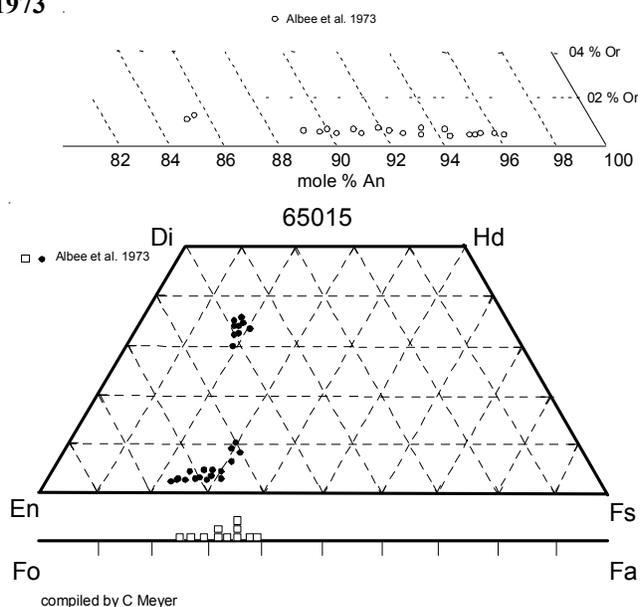


Figure 4: Plagioclase, pyroxene and olivine composition for 65015 (data replotted from Albee et al. 1973).

found to be very calcic ( $An_{93-95}$ ). Laths of plagioclase ( $An_{85-90}$ ) are found in the mesostasis between the oikocrysts (figure 4). Albee et al. (1973) found that the trace FeO content of the plagioclase in the clasts was lower than for the plagioclase in the groundmass. Meyer et al. (1974) found that the trace element (Li, Ba etc) content of the plagioclase phenocrysts was well below the level expected for plagioclase in equilibrium with the bulk rock content. Papanastassiou and

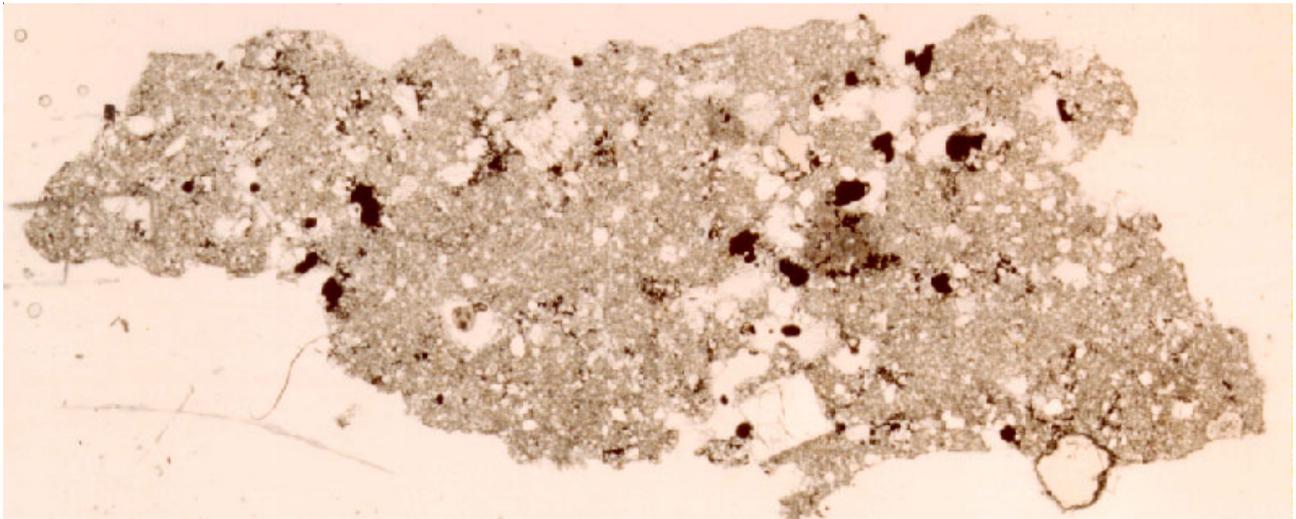


Figure 5: Thin section of 65015. Field of view about 1 inch.

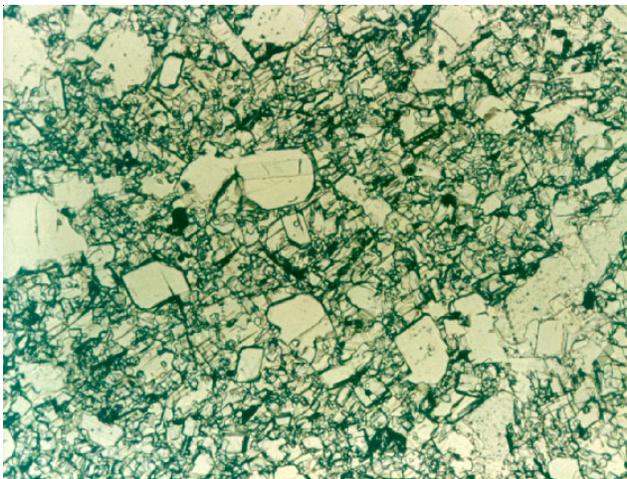


Figure 6: Closeup photomicrograph of thin section of 65015 in plain polarized light. Field of view is about 5 mm.

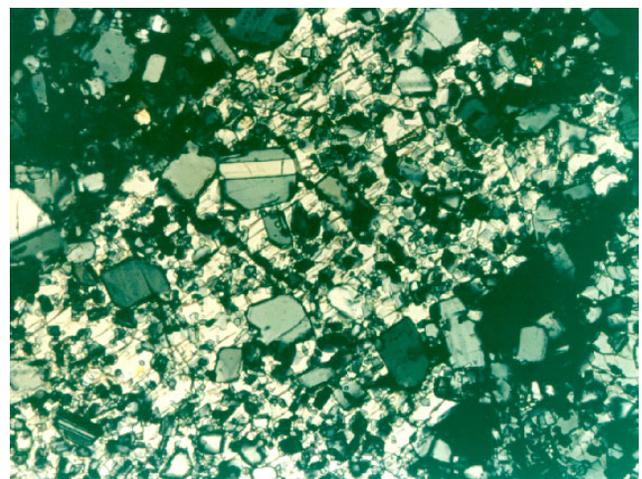


Figure 7: Same area as figure 6, but with crossed polarizers, illustrating large oikocryst of low-Ca pyroxene enclosing numerous chadocrysts of plagioclase.

Wasserburg (1972) also showed that the plagioclase was not in isotopic equilibrium (figure 10).

**Ilmenite:** Ilmenite occurs as irregular poikilitic grains intergrown with other minor phases in the mesostasis (figure 8). It rarely has exsolution of rutile, baddeleyite and Cr-rich spinel (Albee et al. 1973).

**Metal:** El Goresy et al. (1973), Albee et al. (1973) and Misra and Taylor (1975) analyzed the Fe-metal grains (including one large ball, 400 microns) and found they were generally homogeneous 5-6% Ni, 0.2-0.3% Co, 0.12% P and 0.01 % S. Taylor et al. (1976) found that this composition did not change during annealing experiments.

**Whitlockite:** Long needles (40 microns) of whitlockite are the only phosphate (Albee et al. 1973).

### Chemistry

Hubbard et al. (1973), Haskin et al. (1973), Wanke et al. (1977) and others have analyzed 65015 (table 1). The rare earth elements are in high abundance and similar to KREEP (figure 9). The meteoritic siderophiles are relatively high (Ir > 10 ppb) indicating that this is not a pristine rock, but rather an impact melt (Krahenbuhl et al. 1973, Wanke et al. 1977) contaminated with a significant meteoritic component.

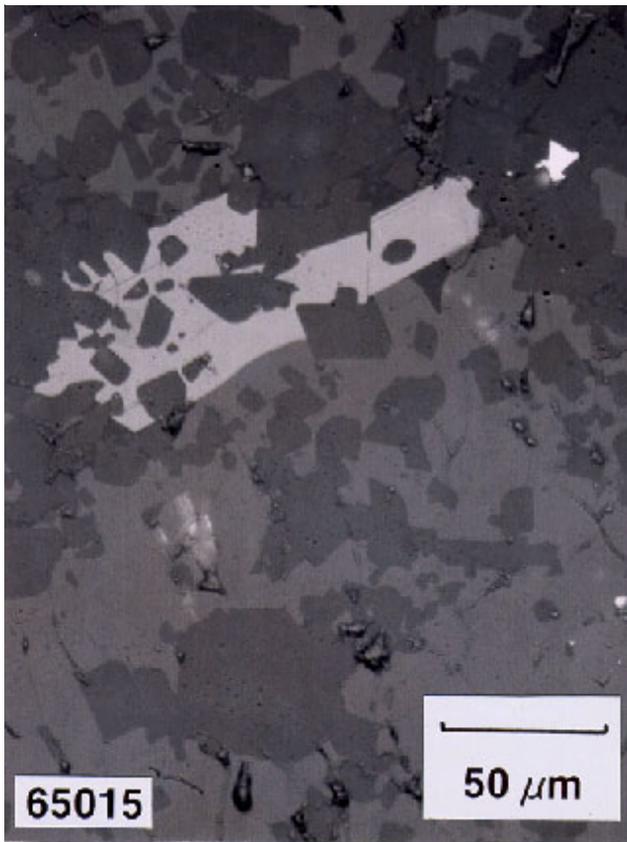


Figure 8: Ilmenite and metal grains in 65015. Reflected light photo of polished thin section.

### **Radiogenic age dating**

Because of the included old component, it has proven difficult to accurately date 65015. Papanastassiou and Wasserburg (1972) found that the plagioclase separates were not on the isochron defined by quintessence and whole rock (figure 10). But these and various other investigators (figure 11) have concluded that 65015 was last melted or “metamorphosed” at about 3.9 b.y. (see table).

Lugmair and Carlson (1978) determined a “model age” of  $4.32 \pm 0.12$  b.y. by Sm-Nd. Nunes et al. (1973) found that U/Pb ages were concordant at 3.99 b.y. and that initial Pb was extremely low. Schaeffer et al. (1978, 1979) studied laser extraction of Ar.

### **Cosmogenic isotopes and exposure ages**

Kirsten et al. (1973) reported a cosmic ray exposure age of  $365 \pm 20$  m.y. for 65015 by the  $^{38}\text{Ar}$  method. Jessberger et al. (1974) found it to be about 490 m.y. Bhandari et al. (1973) determined the “suntan” age as 1.2 m.y. by track studies.

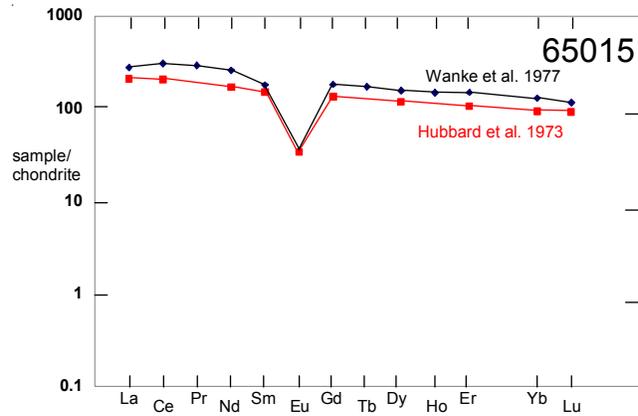


Figure 9: Normalized rare-earth-element diagram for 65015 (data from table 1).

### **Other Studies**

Sato (1976) measured the in-situ oxygen fugacity as a function of temperature.

The magnetic properties were studied by Brecher (1977) and Stephenson et al. (1977); electrical properties were determined by Olhoeft et al. (1973) and Alvarez (1977); thermal expansion by Todd et al. (1973); electron spin resonance by Tsay and Live (1974) and ultraviolet reflectance spectra by Hapke et al. (1978).

### **Processing**

In 1972 a small slab was cut from this rock (figure 12). Allocations were also made from the small butt end.

#### List of Photo #s 65015

S72-39208-39213	Color
S72-43925-43940	B & W (in focus)
S72-47353-47376	slab parts
S76-23580	1 <sup>st</sup> cut
S76-23578	2 <sup>nd</sup> cut

**Table 1a. Chemical composition of 65015.**

<i>reference weight</i>	Hubbard 73	Hubbard 73	Duncan 73	Haskin 73	Krahenbuhl 73	Boynton 75	Wanke 77	Garg 76
						Wasson 77		
SiO <sub>2</sub> %	47.18	(a)	46.99	(a) 47.6	(b)		46.77	
TiO <sub>2</sub>	1.04	(a) 1.1	(d) 1.26	(a) 1.23	(b)	1.2	1.22	1.27
Al <sub>2</sub> O <sub>3</sub>	19.98	(a) 23.2	(c) 19.68	(a) 19.5	(b)	20.03	20.6	19.33
FeO	7.91	(a) 8.71	(b) 8.59	(a) 8.52	(b)	7.98	8.33	8.81
MnO	0.12	(a)	0.112	(a) 0.11		0.13	0.11	0.11
MgO	10.34	(a) 9.5	(b) 9.31	(a) 9.66	(b)		8.34	9.61
CaO	12.03	(a) 12	(b) 11.9	(a) 12.3	(b)	12.73	11.75	11.98
Na <sub>2</sub> O	0.44	(a) 0.59	(b) 0.55	(a) 0.57		0.63	0.52	0.57
K <sub>2</sub> O	0.32	(a) 0.34	(d) 0.36	(a) 0.35	(b)			0.36
P <sub>2</sub> O <sub>5</sub>	0.4	(a)	0.409	(a)				0.39
S %	0.13	(a)	0.079	(a)				0.139
<i>sum</i>	99.89		99.24					99.339
Sc ppm				14.9	(e)	17	16.1	17.1
V							40	36.7
Cr		2350	(d)	1280	(e)	1270	1320	1300
Co				27.1	(e)		49	42.6
Ni			185	(a) 400	(e) 580	(f) 30		730
Cu			1.8	(a)				5.06
Zn					0.48	(f)		1.46
Ga				4.25	(e)			3.81
Ge ppb					380	(f)		600
As								0.17
Se					0.245	(f)		0.29
Rb		9.09	(d) 9.55	(a) 9.2	(e) 7.8	(f)		10.2
Sr		164	(d) 157	(a)				146
Y			179	(a)				174
Zr			909	(a)			920	940
Nb			55.2	(a)				933
Mo								
Ru								
Rh								
Pd ppb								
Ag ppb					1.5	(f)		
Cd ppb					9.25	(f)		
In ppb								
Sn ppb								
Sb ppb					4.34	(f)		
Te ppb					3.25	(f)		
Cs ppm				0.42	(e) 0.405	(f)		0.423
Ba		492	(d) 609			460	560	570
La		49.5	(d)	60	(e)	47	55.6	65.5
Ce		125	(d)	153	(e)	180	132	185
Pr				98	(e)			26
Nd		78	(d)				80	118
Sm		22.2	(d)	26.5	(e)	35	24.8	26.7
Eu		1.91	(d)	2.12	(e)	2.28	2.1	1.97
Gd		26.6	(d)	33	(e)			36
Tb				5.2	(e)	5.4	5.2	6.28
Dy		28.8	(d)	38	(e)		33	38.1
Ho				7.4	(e)			8.3
Er		16.8	(d)	22	(e)			23.8
Tm								
Yb		15.3	(d)	19.1	(e)	20.8	17.1	21.1
Lu		2.24	(d)	2.6	(e)	2.87	2.44	2.84
Hf				19.3	(e)	22	19.2	21.6
Ta						2.4	2.1	2.46
W ppb								883
Re ppb					1.16	(f)		1.8
Os ppb								
Ir ppb					11.6	(f)		17
Pt ppb						(f)		
Au ppb					10.2	(f)		10.9
Th ppm						10.7	9.4	8.92
U ppm		3.42	(d)		3.42	(f)	2.6	2.22

*technique (a) XRF, (b) AA, (c) colorimetry, (d) IDMS, (e) INAA, (f) RNAA*

**Table 1b. Chemical composition of 65015.**

<i>reference weight</i>	Lugmair 78	Nyquist 73	Nunes 73	Ehmann 74	Miller 74	Taylor 73	LSPET 73
SiO <sub>2</sub> %						47.1	(g)
TiO <sub>2</sub>						1.16	(g)
Al <sub>2</sub> O <sub>3</sub>						20.7	(g)
FeO					8.1	(e) 7.37	(g)
MnO							
MgO						9.58	(g)
CaO						11.9	(g)
Na <sub>2</sub> O						0.52	(g)
K <sub>2</sub> O						0.47	(g) 0.48 (h)
P <sub>2</sub> O <sub>5</sub>							
S %							
<i>sum</i>							
Sc ppm					14.4	(e) 9.5	(g)
V						44	(g)
Cr					1230	(e) 1400	(g)
Co					30	(e) 25	(g)
Ni						260	(g)
Cu						6.5	(g)
Zn							
Ga							
Ge ppb							
As							
Se							
Rb	9.09		(d)			8.3	(g)
Sr	163.8		(d)				
Y						182	(g)
Zr				920	1010	(e) 960	(g)
Nb						58	(g)
Mo							
Ru							
Rh							
Pd ppb							
Ag ppb							
Cd ppb							
In ppb							
Sn ppb						400	(g)
Sb ppb							
Te ppb							
Cs ppm						0.35	(g)
Ba						540	(g)
La					56	(e) 58	(g)
Ce						145	(g)
Pr						21	(g)
Nd	101		(d)			81	(g)
Sm	28		(d)			24.5	(g)
Eu					2	(e) 2	(g)
Gd						32	(g)
Tb						4.86	(g)
Dy						30.5	(g)
Ho						7.44	(g)
Er						21	(g)
Tm						3.1	(g)
Yb						19	(g)
Lu						2.9	(g)
Hf				19.8	21.4	(e) 19	(g)
Ta							
W ppb							
Re ppb							
Os ppb							
Ir ppb							
Pt ppb							
Au ppb							
Th ppm			10.32	(d)		10.4	(g) 10 (h)
U ppm			2.58	(d)		2.34	(g) 3 (h)

*technique (a) XRF, (b) AA, (c) colorimetry, (d) IDMS, (e) INAA, (f) RNAA, (g) SSMS, (h) radiation counting*

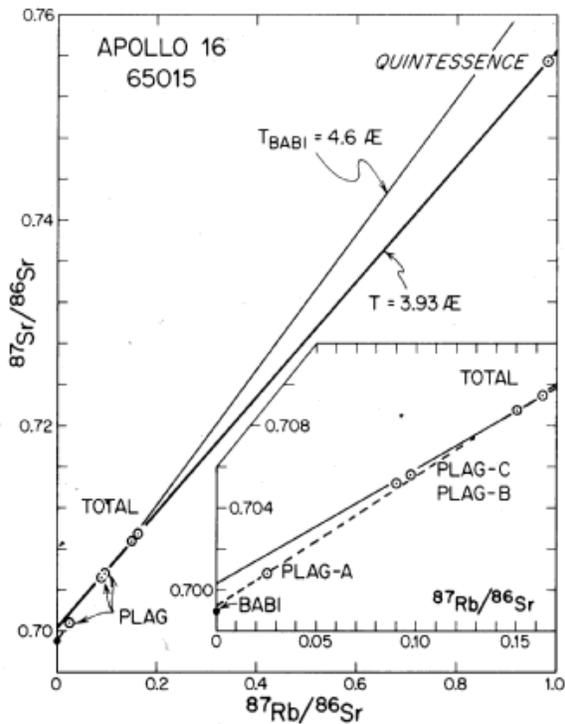


Figure 10: Rb-Sr isochron diagram for 65015 (from Papanastassiou et al. 1972).

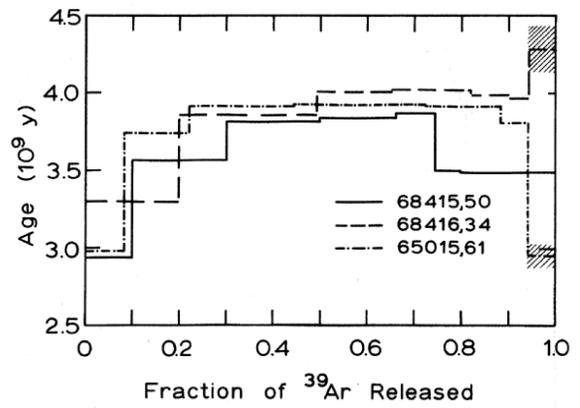
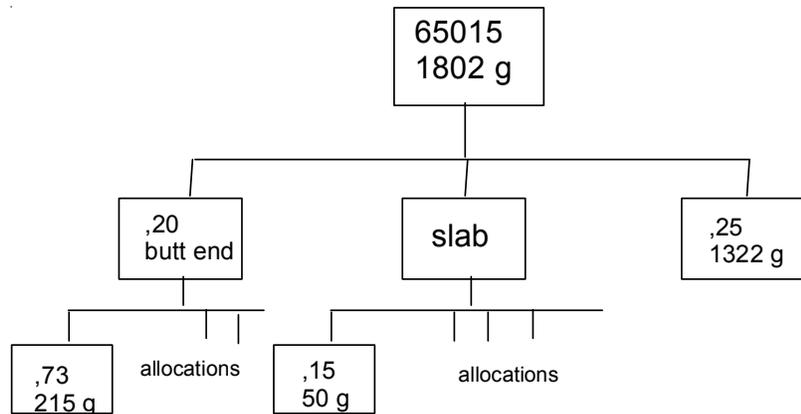
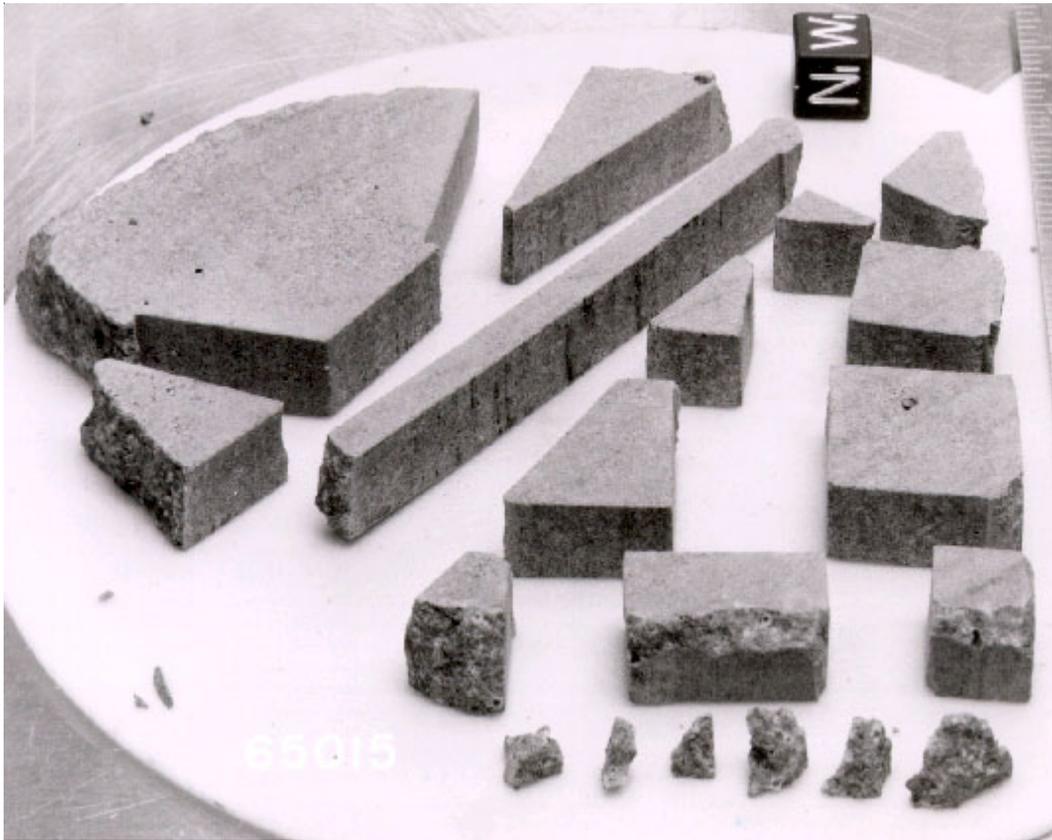


Figure 11: Ar release pattern for 65015 (figure from Kirsten et al. 1973).



### Summary of Age Data for 65015

	Ar-Ar	Rb/Sr	Sm/Nd	Pb/Pb
Papanastassiou and Wasserburg 1972		3.93 ± 0.02 b.y.		
Kirsten et al. 1973	3.92 ± 0.04			
Jessberger et al. 1974	~3.98			
Nunes et al. 1973				3.99



*Figure 12: Group photo of pieces cut from slab of 65015. The cube is 1 cm. NASA photo # S72-47359.*